Learning Propagators for Sea Surface Height Forecasts Using Koopman Autoencoders

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Key Points:

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7	•	We train a neural network to learn a low-dimensional representation of sea sur-
8		face height that facilitates regional predictions
9	•	The approach can work well in situations where linear inverse models struggle, such
10		as on daily-averaged data
11	•	Reconstruction skill highlights sources of predictability, such as the low-latitudes
12		for North Pacific daily sea surface heights

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13 Abstract

- ¹⁴ Due to the wide range of processes impacting the sea surface height (SSH) on daily-to-
- ¹⁵ interannual timescales, SSH forecasts are hampered by numerous sources of uncertainty.
- ¹⁶ While statistical-dynamical methods like Linear Inverse Modeling have been successful

at making forecasts, they often rely on assumptions that can be hard to satisfy given the

- ¹⁸ nonlinear dynamics of the climate. Here, we train convolutional autoencoders with a dy-
- ¹⁹ namical propagator in the latent space to generate forecasts of SSH anomalies. Learn-
- ing a nonlinear dimensionality reduction and the prediction timestepping together results in a propagator that produces better predictions for daily- and monthly-averaged
- 22 SSH in the North Pacific and Atlantic than if the dimensionality reduction and dynam-
- ics are learned separately. The reconstruction skill of the model highlights regions in which
- better representation results in improved predictions: in particular, the tropics for North
- ²⁵ Pacific daily SSH predictions and the Caribbean Current for the North Atlantic.

²⁶ Plain Language Summary

Forecasts of sea surface heights are impacted by numerous sources of uncertainty. 27 While statistical methods for representing temporal changes in the climate system have 28 been useful for making predictions, they often rely on assumptions that do not always 29 hold due to the complex interactions in the climate system. Here, we make a machine 30 learning model that learns a compressed representation of the climate system which fa-31 cilitates sea surface height predictions. The learned compressed representation of the cli-32 mate system results in better sea surface height predictions than would occur if the di-33 mensionality reduction and prediction is done separately. Our machine learning model 34 also points to regions where more accurately representing sea level can result in better 35 regional-scale predictions. 36

37 1 Introduction

The large variety of processes impacting sea surface heights (SSH) on daily-to-interannual 38 timescales implies that forecasts of SSH on these time horizons are hindered by numer-39 ous sources of uncertainty. SSH variability on these timescales is driven by factors in-40 cluding barotropic adjustment to wind stress (Hermans et al., 2022; Kamp et al., 2024; 41 Vinogradova et al., 2007), local air-sea buoyancy fluxes (Cabanes et al., 2006; Gill & Niller, 42 1973), wind-driven Ekman pumping (Webb, 2021; Cabanes et al., 2006), changes in large-43 scale Sverdrup balance (Cabanes et al., 2006), advection of density anomalies (Piecuch 44 & Ponte, 2011), Rossby waves (Chelton & Schlax, 1996; Calafat et al., 2018), buoyancy-45 driven changes in ocean circulation (Roberts et al., 2016), eddy variability due to baro-46 clinic instability (Marques et al., 2022), and the hydrostatic depression of the ocean sur-47 face due to atmospheric pressure anomalies (Piecuch et al., 2016). Developing forecasts 48 for SSH amid these numerous drivers thus presents a challenge. 49

Over the past few decades, statistical-dynamical methods have proven effective for 50 developing forecasts directly from data. Forecasts generated using Linear Inverse Mod-51 els (LIM, Penland (1989); Penland and Sardeshmukh (1995)) have had substantial suc-52 cess in predicting the large-scale evolution of geophysical fields on these timescales (Newman, 53 Shin, & Alexander, 2011; Zanna, 2012; Fraser et al., 2019; Albers & Newman, 2021). The 54 framework generally involves first applying dimensionality reduction to represent the sys-55 tem state using a low-dimensional state vector, and then determining a linear propaga-56 tor using the time-lagged covariance statistics between the state variables. This approach 57 is based on the assumption that the state evolution can be represented as the sum of slow, 58 predictable, linear dynamics and fast, unpredictable, nonlinear dynamics modelled by 59 Gaussian noise (Hasselmann, 1976). Despite the simplicity of such models, LIMs have 60 demonstrated skill comparable to operational forecasting models in some cases (Albers 61 & Newman, 2021; Shin & Newman, 2021; Richter et al., 2020). 62

One appealing aspect of LIMs is the simplified representation of the dynamics as 63 a low-dimensional, linear propagator. While nonlinear dynamical systems can be chaotic, 64 unpredictable, and nontrivial to solve, linear dynamical systems readily admit closed-65 form solutions and can be solved in a systematic manner. The eigenvalues of the prop-66 agator can be used to identify dominant timescales for the dynamics of the system as 67 well as optimal initial conditions for producing anomaly growth (Penland & Sardeshmukh, 68 1995; von Storch et al., 1995; Vimont et al., 2014; Zanna, 2012). However, ensuring that 69 the state evolution is plausibly described by a linear stochastic dynamical system is of-70 ten challenging. Whether or not dynamics can be represented as such depends on the 71 processes being represented and the temporal resolution of the data. The computed prop-72 agator typically depends on the time lag used to compute it, due to nonstationary statis-73 tics (Penland & Sardeshmukh, 1995), unrepresented processes (Penland & Ghil, 1993), 74 fundamental deficiencies in representing dynamical systems using Markov models (DelSole, 75 2000), and sampling of intrinsic oscillatory modes of the system (Penland, 2019). 76

Another sensitivity lies in the application of dimensionality reduction. Clearly, the 77 number of dimensions used to represent the state is a parameter (Newman, Alexander, 78 & Scott, 2011). Additionally, the performance of a LIM may depend on the dimension-79 ality reduction technique applied. Typically, Principal Component Analysis (PCA), also 80 known as Empirical Orthogonal Function analysis in the geosciences, is used to reduce 81 the dimensionality of the system (Hotelling, 1933; Pearson, 1901; Lorenz, 1956). How-82 ever, the requirement that modes are orthogonal can be restrictive (Dommenget & Latif, 83 2002). Alternatively, neural network autoencoders can relax the assumptions of linear-84 ity and orthogonality to obtain more efficient low-dimensional embeddings (Kramer, 1991; 85 Hinton & Salakhutdinov, 2006). Nevertheless, it is unclear whether a more efficient yet 86 complex representation will result in better predictions. 87

Complementing the linear-stochastic dynamical systems framework in inverse mod-88 eling of the earth system is the burgeoning set of data-driven approaches based on the 89 operator-theoretic perspective of nonlinear dynamics. Under Koopman operator theory, 90 nonlinear dynamical systems are represented through the linear (but infinite-dimensional) 91 Koopman operator, which advances measurements of the system through time (Koopman, 92 1931). Thus, obtaining low-dimensional representations of the Koopman operator is a 93 key goal of data-driven dynamical systems modeling. For instance, Dynamic Mode De-94 composition seeks to find the best-fit linear model that advances linear measurements 95 of the system (Schmid, 2010); however, such linear measurements may be insufficient to 96 capture the complexities of nonlinear systems. Therefore, recent deep-learning approaches 97 have modified the autoencoder architecture to learn nonlinear transformations into latent spaces in which the dynamics are approximately linear (Mardt et al., 2018; Lusch 99 et al., 2018; Champion et al., 2019; Yeung et al., 2019; Brunton & Kutz, 2022). 100

Here, we leverage the Koopman Autoencoder framework in Lusch et al. (2018) to construct a linear propagator for SSH prediction on daily-to-interannual timescales in the North Pacific and North Atlantic. We assess the forecasts made by this model relative to baselines in which the dimensionality reduction and propagator are learned separately. We examine the areas of reconstruction skill to interpret how the Koopman Autoencoder attains its performance.

107 2 Methods

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2.1 Data

We use daily- and monthly-averaged simulated SSH fields from the Community Earth System Model, version 2 (CESM2) Large Ensemble dataset (LENS2, Rodgers et al. (2021); Danabasoglu et al. (2021)). The data is from the 250-year simulation period spanning 1850–2100, with radiative forcing following the historical record from 1850–2014 and the CMIP6 SSP3-7.0 forcing scenario thereafter (Danabasoglu et al., 2020; O'Neill et al., 2016).
 Fields are detrended using a locally-fitted fifth-degree polynomial and deseasonalized by

¹¹⁵ removing climatological daily averages.

116 Sea surface heights η are computed by

$$\eta(x, y, t) = \zeta(x, y, t) + \eta_{\rm ib}(x, y, t) \tag{1}$$

where ζ is the dynamic sea level simulated by CESM2 and η_{ib} is the inverse barometer contribution to sea level (Ponte, 2006; Gregory et al., 2019), given by

$$\eta_{\rm ib}(x, y, t) = -\frac{1}{\rho_0 g} p'_a(x, y, t).$$
(2)

Here, $p'_a(x, y, t) = p_a(x, y, t) - \frac{1}{A} \int_A p_a(x, y, t) dA$ is the sea level pressure deviation from the spatial average over the ocean area A at time t, $\rho_0 = 1025$ kg m⁻³ is the reference sea surface density (Smith et al., 2010; Fofonoff & Millard Jr, 1983), and g = 9.81 m s⁻¹ is the acceleration due to gravity.

We use nine ensemble members, with seven members for training and one member for validation and testing. We focus on two regions: the North Pacific $(15^{\circ}S-60^{\circ}N, 115^{\circ}E-60^{\circ}W)$ and the North Atlantic $(5^{\circ}-65^{\circ}N, 60^{\circ}W-0^{\circ}E)$. For training, fields are standardized using the area-weighted mean and standard deviation averaged over all samples in the training set (LeCun et al., 2002). Locations corresponding to land points are masked with zeros.

2.2 Koopman Autoencoder

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Figure 1 illustrates the Koopman Autoencoder (Lusch et al., 2018). The network functions as a propagator for a dynamical system with the entire SSH field as its state variable: it consumes input fields of SSH at a given timestep $n(x_n)$ and outputs the predicted SSH field at the next timestep (\hat{x}_{n+1}) . We use a timestep of one day for networks trained on daily averages and one month for networks trained on monthly averages.

We employ a convolutional architecture that is well-suited for the spatial fields comprising our system state (Fukushima, 1980; LeCun et al., 1989). The encoder E takes in the state vector x_n , extracts features using convolutional filters and transforms the inputs to a lower dimensional embedding z_n . Then, a linear layer L is applied to the latent embedding, functioning as a single propagation timestep. Finally, the decoder Dtransforms the encoded prediction back into the state space, using the state at the next timestep x_{n+1} as the target.

During training, parameters in the Koopman Autoencoder are adjusted through backpropagation (Rumelhart et al., 1986) to optimize a combination of different objective functions in accordance with Lusch et al. (2018).

1. The reconstruction error

$$\mathcal{L}_{\text{reconst}}(x_n) = \|x_n - D(E(x_n))\|_{2,w}^2, \tag{3}$$

where $\|\cdot\|_{2,w}$ is the area-weighted ℓ^2 -norm (see Supporting Text S2). This loss ensures that the encoder and decoder learns a maximally-efficient representation of the SSH in the *d*-dimensional latent space.

2. The prediction error

$$\mathcal{L}_{\text{pred}}(x_n, \dots, x_{n+k}) = \frac{1}{k} \sum_{\ell=1}^k \|x_{n+\ell} - D(L^{\ell} E(x_n))\|_{2,w}^2$$
(4)

The norm $||x_{n+1} - D(LE(x_n))||_{2,w}^2$ indicates the prediction error incurred during a single propagation timestep. In practice, better predictions are obtained by



Figure 1. Koopman Autoencoder schematic. The encoder and decoder are denoted by the brackets labelled E(x) and D(z), respectively, and the inset shows the linear propagator. Yellow blocks indicate convolutional layers, and orange shading indicates ReLU activations. Red blocks indicate pooling layers, and green blocks indicate upsampling layers.

152	penalizing ℓ -timestep predictions for $\ell \in \{1, \ldots, k\}$, where the ℓ -timestep pre-
153	diction $\hat{x}_{n+\ell}$ is given by ℓ applications of the propagator L to the latent embed-
154	ding: $\hat{x}_{n+\ell} = D(L^{\ell}E(x_n))$. We use $k = 20$ recurrent passes for all of our net-
155	works.

¹⁵⁶ We also add a latent space prediction error

$$\mathcal{L}_{\text{linear}}(x_n, x_{n+1}) = \|LE(x_n) - E(x_{n+1})\|_2^2 \tag{5}$$

which further ensures that the linear prediction $\hat{z}_{n+1} = Lz_n = LE(x_n)$ approximates the latent state at the next timestep $z_{n+1} = E(x_{n+1})$. This term may be redundant as our propagator L is not equipped with activations, but is added for consistency with the proposed methodology of Lusch et al. (2018).

¹⁶¹ The net loss is given by

$$\mathcal{L}(x_n, \dots, x_{n+k}) = \lambda_1 \mathcal{L}_{\text{reconst}}(x_n) + \lambda_2 \mathcal{L}_{\text{pred}}(x_n, \dots, x_{n+k}) + \lambda_3 \mathcal{L}_{\text{linear}}(x_n, x_{n+1})$$
(6)

where λ_1 , λ_2 , and λ_3 are hyperparameters. By optimizing this loss, the dimensionality reduction and the timestepping are learned together. This way, the dimensionality reduction is constructed in such a way that predictions are improved.

Separate networks are trained for each region and timescale. Full details about the
 training architecture and procedure are given in Supporting Text S1.

167 **2.3 Baselines**

We contrast the predictions made with our Koopman Autoencoders with baselines in which the dimensionality reduction and predictions are done separately. For dimensionality reduction, we consider Principal Component Analysis (PCA) and Convolutional Autoencoders (CAE). For forecasting, we apply Damped Persistence (DP) and Linear Inverse Modeling (LIM). Prediction baselines are thus determined by combining the two techniques, and are denoted according to the dimensionality technique and propagator used, e.g. "PCA-LIM" or "CAE-DP."

2.3.1 Dimensionality reduction techniques

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As a first baseline, PCA is applied to reduce the dimensionality of the state. In PCA, the data is linearly projected onto the *d*-dimensional subspace that maximizes the variance of the data. As a result, dimensions describing the data are linear and orthogonal, a restriction that may result in poor representation of nonlinear data manifolds.

As a nonlinear alternative to PCA, we also train Convolutional Autoencoders (CAE). Autoencoders generalize PCA by allowing for nonlinear transformations to a latent space and can learn more efficient representations than PCA (Kramer, 1991; Hinton & Salakhutdinov, 2006; Shamekh et al., 2023; Oommen et al., 2022). For the CAE, we use an encoder and decoder with the same architectures as those of the Koopman Autoencoder, and we train it with nearly identical hyperparameters (see Supporting Text S1).

186 2.3.2 Predictions in the latent space

¹⁸⁷ We compare the forecasts made by the Koopman Autoencoder to Damped Persis-¹⁸⁸ tence (DP, Lorenz (1973)). Given a latent state z_n , the prediction at lag τ is given by

$$\hat{z}_{n+\tau} = \mathbf{D}(\tau) z_n \tag{7}$$

where $\mathbf{D}(\tau)$ is a diagonal matrix whose entries give the autocorrelation of each of the latent variables at lag τ . The propagator $\mathbf{D}(\tau)$ is computed iteratively for each time lag by first selecting a training timescale τ_0 , computing the lag- τ_0 autocorrelations $\mathbf{D}_0 =$ $\mathbf{D}(\tau_0)$, and then defining $\mathbf{D}(\tau) = (\mathbf{D}_0)^{\tau/\tau_0}$. For a fair comparison with the Koopman Autoencoder, we set τ_0 by fitting DP models using $\tau_0 \in \{1, \ldots, k\}$ and selecting the model with the lowest average prediction error on timesteps 1 to k on the validation dataset.

We also explore predictions made by a Linear Inverse Model (LIM, Penland (1989)). The underlying assumption behind LIM is that the dynamics of a system can be wellrepresented as a linear dynamical system forced by noise:

$$\frac{dz}{dt} = \mathbf{A}z + \xi \tag{8}$$

where ξ is sampled from a Normal distribution. Then, the evolution matrix **A** can be estimated through an error minimization procedure as

$$\mathbf{A} = \frac{1}{\tau_0} \log \left(\mathbf{C}(\tau_0) \mathbf{C}(0)^{-1} \right) \tag{9}$$

where $\mathbf{C}(\tau) = \langle z(t+\tau)z^T(t) \rangle$ gives the time- τ lagged covariance (with angled brackets denoting a time average) and τ_0 is a fitted timescale. Predictions are then given by

$$\hat{z}_{n+\tau} = \mathbf{B}(\tau) z_n \tag{10}$$

with the propagator $\mathbf{B}(\tau)$ given by

$$\mathbf{B}(\tau) = \exp(\mathbf{A}\tau) = \exp\left[\frac{\tau}{\tau_0}\log\left(\mathbf{C}(\tau_0)\mathbf{C}(0)^{-1}\right)\right]$$
(11)

The covariance matrix is computed over all ensemble members, and again τ_0 is selected

timesteps 1 through k.

by fitting LIMs for $\tau_0 \in \{1, \dots, k\}$ and selecting the model with the lowest error over

In order for a LIM to be valid, several conditions should be met. One basic crite-206 rion is that the learned propagator should be stable with decaying eigenvalues. (Simi-207 larly, the eigenvalues of the propagator learned by the Koopman Autoencoder should also 208 decay.) Supporting Figure S1 verifies that all propagators considered in this study are stable. Another requirement is that the evolution matrix defined by Equation 9 must 210 be independent of the time lag τ_0 used to compute it. However, this is a strong crite-211 rion to meet; common practice is to compute the matrix norm of the propagator $||A||_2$ 212 for different τ_0 and to select a propagator based on a timescale τ_0 in which the matrix 213 norm is relatively constant. Supporting Figure S2 shows the ℓ^2 -matrix norms of the evo-214 lution matrix of the LIM baselines on the range $\tau_0 \in \{1, \ldots, k\}$; over this range, the 215 matrix norm varies by over 300% for all of the regions and timescales considered. 216

217 **3 Results**

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In this section, we compare the forecasts made by the Koopman Autoencoder to the other baselines. We use the Mean Square Error (MSE) and Pattern Correlation Coefficient (Legates & Davis, 1997) to assess our predictions, as well as MSE-based skill scores (Murphy, 1988). Metrics are defined explicitly in Supporting Text S2.

3.1 Evaluating prediction performance

Figure 2 compares the area-weighted prediction MSE and pattern correlation of 223 SSH predictions of the Koopman Autoencoder to the baselines using d = 20 latent di-224 mensions on forecast lead times τ of up to $\tau_{max} = 120$ days (daily data) and $\tau_{max} =$ 225 36 months (monthly data). The CAE generally has the lowest reconstruction error for 226 all dimensionality reduction techniques, beating PCA MSE by a margin of 2-4% at lead 227 $\tau = 0$ for all regions and timescales except in the North Atlantic on monthly data (See 228 Supporting Table S1). The Koopman Autoencoder has the worst reconstructions of all 229 the methods considered: over all regions and timescales, MSE is on average 32% higher 230 for the Koopman Autoencoder than for PCA. However, the better reconstruction error 231 of the CAE does not necessarily result in better predictions. In fact, predictions made 232 by applying propagators to CAE modes are often worse than predictions made using PCA 233 for dimensionality reduction (e.g., using a DP propagator for North Pacific daily SSH, 234 Figure 2a). In contrast, the Koopman Autoencoder generally results in better predic-235 tions than the baselines as measured by the area-weighted MSE and pattern correlation. 236 Supporting Table S2 quantitatively summarizes the forecast performance of the mod-237 els in Figure 2 through the skill score of the different prediction methods relative to PCA-238 DP, averaged over forecast leads up to $\tau_{\rm max}.$ Skill of the models relative to PCA-DP de-239 pends significantly on the region and timescales considered but averaged over all regions 240 and timescales, PCA-LIM has about 6.8% skill over PCA-DP, skill of CAE-LIM is slightly 241 worse than PCA-LIM (6.4%), and skill of the Koopman Autoencoder is the highest (8.4%). 242 In effect, by learning the dynamics and the dimensionality reduction together, the Koop-243 man Autoencoder learns a nonlinear latent-space representation of the state that implic-244 itly results in better SSH predictions. 245

The advantages of using the Koopman Autoencoder over, for example, PCA-LIM 246 are more apparent on daily timescales than on monthly timescales. In the North Pacific, 247 prediction skill of the Koopman Autoencoder relative to PCA-DP on daily-averaged data 248 is 4.5% higher than that of PCA-LIM but is only 3.0% higher for monthly-averaged data; 249 in the North Atlantic, Koopman skill is 1.1% higher than PCA-LIM on daily data but 250 is 1.3% lower on monthly data. One potential reason is that the assumptions underly-251 ing LIM may be better satisfied for monthly averages than daily averages, because monthly-252 averaged fields smooth out small-scale, nonlinear features (Sardeshmukh & Sura, 2009; 253 Stephenson et al., 2004). The Koopman Autoencoders also outperform PCA-LIM by a 254 wider margin in the North Pacific than in the North Atlantic. This may be due to the 255

fact that the inverse barometer component constitutes a larger share of the SSH variability in the North Atlantic region considered (about 71% in the North Atlantic on daily timescales vs 32% in the North Pacific; see Supporting Figure S3). This high-frequency variability may be well-represented by white noise, again underpinning the relative suc-

260 cess of PCA-LIM.



Figure 2. Forecast MSE and Pattern Correlation in the North Pacific and North Atlantic on daily and monthly timescales. Colors indicate dimensionality reduction techniques (red for the Koopman Autoencoder, blue for PCA, and light green for CAE), while markers indicate propagation techniques (x's for the Koopman Autoencoder, filled circles for LIM, and open circles for DP). The black dotted line indicates the climatological MSE of SSH.

3.2 Sensitivity to the number of dimensions

Both the dimensionality reduction and learned propagator's predictions may be sensitive to the dimensionality of the latent space. Figure 3 explores both of these sensitivities. Due to the computational cost of training each network, sensitivity is examined only in one region and timescale; we focus on forecasts of daily-averaged SSH in the North Pacific as the Koopman Autoencoder was shown to generate skillful predictions for these dynamics.

As shown in Figure 3a and Supporting Table S3, reconstruction performance improves as the number of latent dimensions is increased up to d = 40 for all dimensionality reduction techniques considered. Just as in Section 3.1, for any given number of dimensions, the CAE has the best reconstructions, outperforming PCA by 2–4%, while the Koopman Autoencoder has the worst reconstructions, with reconstruction MSE 1– 13% higher than that of PCA.

Like the reconstruction skill, the predictions of the Koopman Autoencoder also im-274 prove as the dimensionality is increased, as shown in Figure 3b. Koopman operator the-275 ory suggests that this should be the case, as it states that infinitely many observables 276 must be prescribed to guarantee a nonlinear dynamical system is fully determined. Nev-277 ertheless, the utility of using the Koopman Autoencoder for building propagators dimin-278 ishes as the number of dimensions is increased. Figure 3c shows the domain-averaged 279 prediction skill of the Koopman Autoencoder relative to PCA-LIM predictions using the 280 same dimensionality. For all dimensionalities, the Koopman Autoencoder outperforms 281 PCA-LIM forecasts up to $\tau = 120$ days; however, up to forecast leads of $\tau = 60$ days, 282 the skill of the Koopman Autoencoder decreases as the dimensionality increases. Much 283 of this seems to be simply because the Koopman Autoencoder becomes worse at recon-284 structions relative to PCA for higher latent dimensionalities (e.g., 1% higher MSE for 285 d = 10 vs 13% higher MSE for d = 40; see Supporting Table S3). This suggests that 286 the Koopman Autoencoder approach may be most useful for developing low-dimensional 287 propagators. 288

3.3 Regions of skill

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To understand how the Koopman Autoencoder attains its performance, Figure 4 shows the MSE-skill score of the Koopman Autoencoder relative to PCA-based propagators for daily SSH forecasts in the North Pacific and North Atlantic. We focus on PCAbased propagators because of the simplicity and interpretability of linear, orthogonal dimensionality reduction, which the CAE cannot afford. For example, due to the orthogonality of modes, applying damped persistence to the principal components results in purely local dampening of SSH at each location.

Figure 4a shows domain-averaged MSE skill scores for the Koopman Autoencoder and PCA-LIM relative to PCA-DP. Skill scores for the Koopman Autoencoder and PCA-LIM relative to PCA-DP start at about 0, increase to a maximum at a lead of about 30 days, and gradually taper for longer-term forecasts. However, the Koopman Autoencoder skill is much higher than that of PCA-LIM at all lags—by 72% at lead 5 days and by at least 47% for leads up to 120 days.

Figures 4c-e and 4f-h show the regional variations of Koopman Autoencoder skill 303 relative to PCA-DP and to PCA-LIM, respectively, for a few different lead times. No-304 tably, the Koopman Autoencoder is better at reconstructing SSH than PCA at low lat-305 itudes but is worse at midlatitudes (Figure 4c). However, by lag $\tau = 5$ days, the neg-306 ative skill in the midlatitudes has diminished compared to PCA-LIM (Figure 4g), and 307 there is positive skill relative to PCA-DP over the entire domain (Figure 4d). Because 308 the midlatitude SSH variability is dominated by the high-frequency inverse-barometer 309 component (Supporting Figure S3), midlatitude SSH dynamics are inherently less pre-310



KAE performance by dimensionality, Pacific daily SSH

Figure 3. Sensitivity of the Koopman Autoencoder to number of dimensions for predicting North Pacific daily-averaged SSH. (a) Reconstruction error by dimensionality for PCA (blue), CAE (light green), and the Koopman Autoencoder (red). (b) Domain-averaged MSE skill scores of the Koopman Autoencoder predictions relative to climatology for different latent space dimensionalities. (c) Domain-averaged skill score of the Koopman Autoencoder relative to equivalent dimensionality PCA-LIM as a function of forecast lead.

dictable than low-latitude dynamics. Therefore, for North Pacific regional-scale predictions, quality representations of SSH in the tropics are much more helpful for regionalscale predictions than representations in the midlatitudes. Because the dimensionality reduction and propagation are learned together in the Koopman Autoencoder, it can deploy its latent dimensions to focus on representing low-latitude SSH initial conditions particularly well. In contrast, when the dimensionality reduction is done separately, dimensions may be wasted on characterizing variability that is not predictable.

The skill maps also highlight dynamics that the PCA-based propagators do not fully 318 capture. For instance, since PCA-DP characterizes the *local* predictability of SSH, skill 319 of the Koopman Autoencoder relative to PCA-DP indicates that it is capturing *nonlo*-320 cal drivers of SSH. Midlatitude skill in the Northeastern Pacific at leads of $\tau = 5$ days 321 (Figure 4d) could come from the advection of sea level pressure anomalies via midlat-322 itude Westerlies, which traverse the Pacific basin on $\mathcal{O}(5-10 \text{ days})$. In the low latitudes, 323 the skill of the Koopman Autoencoder with respect to PCA-DP and PCA-LIM increases 324 until about 30 days (Figures 4a), with the strongest skill occurring in narrow, zonal bands 325 adjacent to the equator (Figure 4h). This timescale and region of enhanced skill is con-326 sistent with the timescale and westward propagation of Equatorial Rossby waves. 327

In the North Atlantic, we note that reconstruction errors for the Koopman Autoencoder at time $\tau = 0$ are poor, with a domain-average skill of -0.14 relative to the PCA reconstructions. However, once again, the latent space representation of the state results in better skill at nonzero time lags up to $\tau = 100$ days (Figure 4b). Figures 4i-k show that the prediction skill of the Koopman Autoencoder occurs primarily in the Atlantic Subtropical Gyre and Gulf Stream separation. Because gyre dynamics are associated primarily with low-variability geostrophic balance, such variability may be underrepresented in variance-targeting PCA-based reconstructions, even though this variability may be predictable on the daily-to-seasonal timescale. Reconstruction skill relative to PCA suggests that the Caribbean Current may be a source of this gyre predictability for SSH predictions in the North Atlantic (Figure 4i).

4 Discussion

Statistical-dynamical models—and linear inverse models in particular—have be-340 come indispensable forecasting tools in the past few decades, owing to their simplicity, 341 interpretability, and skill (Penland & Sardeshmukh, 1995; Alexander et al., 2008; von 342 Storch et al., 1995). Modern techniques can help extract more information from data 343 for nonlinear systems. In this study, we trained convolutional neural networks with em-344 bedded time-stepping to learn a low-dimensional latent space that facilitates predictions 345 of SSH. Training the network to learn the dimensionality reduction and propagation si-346 multaneously tends to result in better forecasts than if the reduction and propagation 347 are learned separately, as done typically with LIM for example. 348

We examined some sensitivities of the Koopman Autoencoder method compared 349 to LIM. The skillfulness of the Koopman Autoencoder is most apparent in situations when 350 the assumptions for LIM are least valid (such as on daily data, where the state vector 351 includes highly nonlinear, small-scale features). Additionally, we examined the sensitiv-352 ity to the dimensionality of the latent space. Our results suggest that the Koopman Au-353 to encoder framework is best for building low-dimensional propagators; however, com-354 putational considerations led us to consider only one region and timescale and up to 40 355 latent dimensions, so the robustness of this result to different dynamics and a wider range 356 of dimensionalities should be further investigated. 357

Spatial variations in the reconstruction skill of the Koopman Autoencoder point 358 to sources of predictability that the Koopman Autoencoder leverages to make better pre-359 dictions than LIM. We identified tropical Pacific SSH as a source of predictability for 360 North Pacific daily-averaged SSH and the Caribbean Current SSH for North Atlantic 361 SSH. One limitation of this study is that a univariate field variable is used for SSH pre-362 dictions. Previous studies have demonstrated that including multiple variables can im-363 prove LIM predictions (Newman, Alexander, & Scott, 2011; Capotondi et al., 2022; Bren-364 nan et al., 2023). Using multiple input channels to incorporate different fields may im-365 prove the Koopman Autoencoder's SSH predictions and reveal additional sources of pre-366 dictability. 367

The focus of this study has been to develop an efficient propagator for SSH and to assess its forecasting skill. The imposed linearity of the dynamics in the latent space could be relaxed (for instance, to obtain better predictions). However, the comprehensive theory underpinning linear systems makes the linear propagator potentially appealing for interpretation, yielding possible advantages in applications like predictability (Vimont et al., 2014; Tziperman et al., 2008), emulation (Beucler et al., 2021; Bi et al., 2023), and inference (Baldovin et al., 2020; Falasca et al., 2024).

One question is how the latent state can be physically interpreted (Shamekh et al., 2023; Behrens et al., 2022). In the context of Koopman operator theory, the latent space variables are nonlinear observables of the dynamical system state, but the nonlinearities in the encoder and decoder make it challenging to interpret what these observables measure. One approach to gaining physical understanding of the latent space could be to probe the sensitivity of the decoder to changes in the latent space, either through observing the sensitivity of the outputs to perturbations to the latent space variables (Oring



Figure 4. Koopman Autoencoder MSE skill scores for daily-averaged North Pacific (a, c–h) and North Atlantic (b, i–n) SSH predictions. (a, b): Domain-averaged skill as a function of lead time. Red: Skill of Koopman Autoencoder relative to PCA-DP. Purple: Koopman Autoencoder relative to PCA+LIM. Cyan: Skill of PCA-LIM relative to PCA-DP. Black dotted lines indicate forecast leads used for panels c–h. (c, d, e, i, j, k): Skill scores of Koopman Autoencoder relative to PCA-DP at select time lags. (f, g, h, l, m, n): Same but for skill relative to PCA-LIM.

et al., 2021; Leeb et al., 2022) or examining the gradients of the decoder (Mamalakis et

al., 2022; Baehrens et al., 2010). Such methods for interpreting the latent space, cou-

 $_{384}$ pled with eigenanalysis for understanding the timescales for the propagator, could help

elucidate the physical processes represented in the latent space, and is left for future work.

Nevertheless, we believe this study has demonstrated a potentially useful approach for developing efficient, low-dimensional linear propagators for climate fields.

388 Appendix A Open Research

The CESM2 Large Ensemble Dataset is available from the NCAR Climate Data 389 Gateway at https://doi.org/10.26024/kgmp-c556 (Danabasoglu et al., 2021). The 390 code used for data processing, training, analysis and visualization in this study, as well 391 as the files for reproducing the software environment, are provided under the MIT license 392 at https://github.com/andrewbrettin/koopman_autoencoders_ssh_prediction (Brettin, 393 2024). Figure 1 was built using the PlotNeuralNet software preserved at https://doi 394 .org/10.5281/zenodo.2526396, which is available via the MIT license (HarisIqbal88, 395 2018). 396

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406 References

407	Albers, J. R., & Newman, M. (2021). Subseasonal predictability of the North At-
408	lantic Oscillation. Environmental Research Letters, 16(4), 044024.
409	Alexander, M. A., Matrosova, L., Penland, C., Scott, J. D., & Chang, P. (2008).
410	Forecasting Pacific SSTs: Linear inverse model predictions of the PDO. Jour-
411	nal of Climate, 21(2), 385–402.
412	Baehrens, D., Schroeter, T., Harmeling, S., Kawanabe, M., Hansen, K., & Müller,
413	KR. (2010). How to Explain Individual Classification Decisions. Journal of
414	Machine Learning Research, 11, 1803–1831.
415	Baldovin, M., Cecconi, F., & Vulpiani, A. (2020). Understanding causation via cor-
416	relations and linear response theory. Physical Review Research, $2(4)$, 043436.
417	Behrens, G., Beucler, T., Gentine, P., Iglesias-Suarez, F., Pritchard, M., & Eyring,
418	V. (2022). Non-Linear Dimensionality Reduction With a Variational Encoder
419	Decoder to Understand Convective Processes in Climate Models. Journal of
420	Advances in Modeling Earth Systems, 14(8), e2022MS003130.
421	Beucler, T., Pritchard, M., Rasp, S., Ott, J., Baldi, P., & Gentine, P. (2021). En-
422	forcing Analytic Constraints in Neural Networks Emulating Physical Systems.
423	Physical Review Letters, $126(9)$, 098302 .
424	Bi, K., Xie, L., Zhang, H., Chen, X., Gu, X., & Tian, Q. (2023). Accurate medium-
425	range global weather forecasting with 3D neural networks. Nature, $619(7970)$,
426	533 - 538.
427	Brennan, M. K., Hakim, G. J., & Blanchard-Wrigglesworth, E. (2023). Monthly Arc-
428	tic Sea-Ice Prediction With a Linear Inverse Model. Geophysical Research Let-
429	ters, 50(7), e2022GL101656.
430	Brettin, A. (2024). Code for "Learning Propagators for Sea Surface Height Forecasts
431	Using Koopman Autoencoders" [Software]. Zenodo. Retrieved from https://
432	
	github.com/andrewbrettin/koopman_autoencoders_ssh_prediction
433	github.com/andrewbrettin/koopman_autoencoders_ssh_prediction Brunton, S. L., & Kutz, J. N. (2022). Data-Driven Science and Engineering: Ma-

435	Press.
436	Cabanes, C., Huck, T., & Colin de Verdière, A. (2006). Contributions of Wind Forc-
437	ing and Surface Heating to Interannual Sea Level Variations in the Atlantic
438	Ocean. Journal of Physical Oceanography, 36(9), 1739–1750.
439	Calafat, F. M., Wahl, T., Lindsten, F., Williams, J., & Frajka-Williams, E. (2018).
440	Coherent modulation of the sea-level annual cycle in the United States by
441	Atlantic Rossby waves. Nature communications, $9(1)$, 2571.
442	Capotondi, A., Newman, M., Xu, T., & Di Lorenzo, E. (2022). An Optimal Precur-
443	sor of Northeast Pacific Marine Heatwaves and Central Pacific El Niño Events.
444	Geophysical Research Letters, 49(5), e2021GL097350.
445	Champion, K., Lusch, B., Kutz, J. N., & Brunton, S. L. (2019). Data-driven dis-
446	covery of coordinates and governing equations. Proceedings of the National
447	Academy of Sciences, 116(45), 22445–22451.
448	Chelton, D. B., & Schlax, M. G. (1996). Global Observations of Oceanic Rossby
449	Waves. Science, 272(5259), 234–238.
450	CISL. (2023). Derecho: HPE Cray EX System (University Community Computing).
451	Boulder, CO: NSF National Center for Atmospheric Research. doi: https://doi
452	.org/10.5065/qx9a-pg09
453	Danabasoglu, G., Deser, C., Rodgers, K., & Timmermann, A. (2021). CESM2
454	Large Ensemble Dataset [dataset]. National Center for Atmospheric Re-
455	search. Retrieved from https://www.earthsystemgrid.org/dataset/
456	ucar.cgd.cesm2le.output.html doi: https://doi.org/10.26024/kgmp-c556
457	Danabasoglu, G., Lamarque, JF., Bacmeister, J., Bailey, D., DuVivier, A., Ed-
458	wards, J., others (2020). The Community Earth System Model Ver-
459	sion 2 (CESM2). Journal of Advances in Modeling Earth Systems, $12(2)$,
460	e2019MS001916.
461	DelSole, T. (2000). A Fundamental Limitation of Markov Models. Journal of the At-
462	$mospheric\ Sciences,\ 57(13),\ 2158-2168.$
463	Dommenget, D., & Latif, M. (2002). A Cautionary Note on the Interpretation of
464	EOFs. Journal of Climate, $15(2)$, $216-225$.
465	Falasca, F., Perezhogin, P., & Zanna, L. (2024). Data-driven dimensionality reduc-
466	tion and causal inference for spatiotemporal climate fields. Physical Review E ,
467	109(4), 044202.
468	Fofonoff, N. P., & Millard Jr, R. (1983). Algorithms for the Computation of Funda-
469	mental Properties of Seawater. UNESCO Technical Papers in Marine Sciences,
470	44.
471	Fraser, R., Palmer, M., Roberts, C., Wilson, C., Copsey, D., & Zanna, L. (2019). In-
472	vestigating the predictability of North Atlantic sea surface height. Climate Dy-
473	namics, 53, 2175-2195.
474	Fukushima, K. (1980). Neocognitron: A Self-organizing Neural Network Model for
475	a Mechanism of Pattern Recognition Unaffected by Shift in Position. <i>Biological</i>
476	Cybernetics, 36(4), 193-202.
477	Gill, A., & Niller, P. (1973). The theory of the seasonal variability in the ocean.
478	Deep Sea Research and Oceanographic Abstracts, 20(2), 141–177.
479	Gregory, J. M., Griffies, S. M., Hughes, C. W., Lowe, J. A., Church, J. A., Fukimori,
480	I., others (2019). Concepts and terminology for sea level: Mean, variability
481	and change, both local and global. Surveys in Geophysics, 40, 1251–1289.
482	Harislqbal88. (2018). <i>PlotNeuralNet</i> [Software]. Zenodo. doi: https://doi.org/10
483	.5281/zenodo.2526396
484	Hasselmann, K. (1976). Stochastic climate models: Part I. Theory. <i>Tellus</i> , 28(6),
485	
486	Hermans, T. H., Katsman, C. A., Camargo, C. M., Garner, G. G., Kopp, R. E.,
487	& Slangen, A. B. (2022). The Effect of Wind Stress on Seasonal Sea-Level
488	Unange on the Northwestern European Shelf. Journal of Climate, $35(6)$, 1745, 1750
489	1(40-1759.

Hinton, G. E., & Salakhutdinov, R. R. (2006). Reducing the Dimensionality of Data 490 with Neural Networks. Science, 313(5786), 504–507. 491 Analysis of a complex of statistical variables into principal Hotelling, H. (1933).492 components. Journal of Educational Psychology, 24(6), 417. 493 Kamp, W., Han, W., Zhang, L., Kido, S., & McCreary, J. P. (2024). Tropical At-494 mospheric Intraseasonal Oscillations Leading to Sea Level Extremes in Coastal 495 Indonesia during Recent Decades. Journal of Climate, 37(9), 2867–2880. 496 Koopman, B. O. (1931). Hamiltonian systems and transformation in Hilbert space. 497 Proceedings of the National Academy of Sciences, 17(5), 315–318. 498 Kramer, M. A. (1991). Nonlinear principal component analysis using autoassociative 499 neural networks. AIChE journal, 37(2), 233–243. 500 LeCun, Y., Boser, B., Denker, J. S., Henderson, D., Howard, R. E., Hubbard, W., 501 & Jackel, L. D. (1989).Backpropagation applied to handwritten zip code 502 recognition. Neural Computation, 1(4), 541-551. 503 LeCun, Y., Bottou, L., Orr, G. B., & Müller, K.-R. (2002). Efficient BackProp. In 504 G. M. et al. (Ed.), Neural Networks: Tricks of the Trade (pp. 9-48). Springer. 505 Leeb, F., Bauer, S., Besserve, M., & Schölkopf, B. (2022).Exploring the Latent 506 Space of Autoencoders with Interventional Assays. Advances in Neural Infor-507 mation Processing Systems, 35, 21562–21574. 508 Legates, D. R., & Davis, R. E. (1997). The continuing search for an anthropogenic 509 climate change signal: Limitations of correlation-based approaches. Geophysi-510 cal Research Letters, 24(18), 2319–2322. 511 Lorenz, E. N. (1956). Empirical Orthogonal Functions and Statistical Weather Pre-512 diction (Vol. 1). Massachusetts Institute of Technology, Department of Meteo-513 rology Cambridge. 514 Lorenz, E. N. (1973). On the Existence of Extended Range Predictability. Journal 515 of Applied Meteorology, 543–546. 516 Lusch, B., Kutz, J. N., & Brunton, S. L. (2018). Deep learning for universal linear 517 embeddings of nonlinear dynamics. Nature Communications, 9(1), 4950. 518 Mamalakis, A., Ebert-Uphoff, I., & Barnes, E. A. (2022). Neural network attribution 519 methods for problems in geoscience: A novel synthetic benchmark dataset. En-520 vironmental Data Science, 1, e8. 521 Mardt, A., Pasquali, L., Wu, H., & Noé, F. (2018). VAMPnets for deep learning of 522 molecular kinetics. Nature communications, 9(1), 5. 523 Marques, G. M., Loose, N., Yankovsky, E., Steinberg, J. M., Chang, C.-Y., Bhamidi-524 NeverWorld2: An idealized model hierarchy to pati, N., ... others (2022).525 investigate ocean mesoscale eddies across resolutions. Geoscientific Model 526 Development, 15(17), 6567–6579. 527 Murphy, A. H. (1988). Skill Scores Based on the Mean Square Error and their Re-528 lationships to the Correlation Coefficient. Monthly Weather Review, 116(12), 529 2417 - 2424.530 Newman, M., Alexander, M. A., & Scott, J. D. (2011). An empirical model of tropi-531 cal ocean dynamics. Climate Dynamics, 37, 1823–1841. 532 Newman, M., Shin, S.-I., & Alexander, M. A. (2011). Natural Variation in ENSO 533 Flavors. Geophysical Research Letters, 38(14). 534 O'Neill, B. C., Tebaldi, C., van Vuuren, D. P., Eyring, V., Friedlingstein, P., Hurtt, 535 G., ... Sanderson, B. M. (2016). The Scenario Model Intercomparison Project 536 (ScenarioMIP) for CMIP6. Geoscientific Model Development, 9(9), 3461–3482. 537 Retrieved from https://gmd.copernicus.org/articles/9/3461/2016/ doi: 538 10.5194/gmd-9-3461-2016 539 Oommen, V., Shukla, K., Goswami, S., Dingreville, R., & Karniadakis, G. E. (2022). 540 Learning two-phase microstructure evolution using neural operators and au-541 to encoder architectures. npj Computational Materials, $\mathcal{S}(1)$, 190. 542 Oring, A., Yakhini, Z., & Hel-Or, Y. (2021, 18–24 Jul). Autoencoder Image Inter-543 polation by Shaping the Latent Space. In Proceedings of the 38th International 544

545	Conference on Machine Learning (Vol. 139, pp. 8281–8290).
546	Pearson, K. (1901). On Lines and Planes of Closest Fit to Systems of Points in
547	Space. The London, Edinburgh, and Dublin Philosophical Magazine and Jour-
548	nal of Science, 2(11), 559–572.
549	Penland, C. (1989). Random Forcing and Forecasting Using Principal Oscillation
550	Pattern Analysis. Monthly Weather Review, 117(10), 2165–2185.
551	Penland, C. (2019). The Nyquist Issue in Linear Inverse Modeling. Monthly Weather
552	Review, 147(4), 1341–1349.
553	Penland, C., & Ghil, M. (1993). Forecasting Northern Hemisphere 700-mb Geopo-
554	tential Height Anomalies Using Empirical Normal Modes. Monthly Weather
555	Review, 121(8), 2355–2372.
556	Penland, C., & Sardeshmukh, P. D. (1995). The optimal growth of tropical sea sur-
557	face temperature anomalies. Journal of Climate, 8(8), 1999–2024.
558	Piecuch, C., Dangendorf, S., Ponte, R. M., & Marcos, M. (2016). Annual sea level
559	changes on the north american northeast coast: Influence of local winds and
560	barotropic motions. Journal of Climate, 29(13), 4801–4816.
561	Piecuch, C., & Ponte, R. (2011). Mechanisms of Interannual Steric Sea Level Vari-
562	ability. Geophysical Research Letters, 38(15).
563	Ponte, R. M. (2006). Low-frequency sea level variability and the inverted barometer
564	effect. Journal of Atmospheric and Oceanic Technology, 23(4), 619–629.
565	Richter, I., Chang, P., & Liu, X. (2020). Impact of Systematic GCM Errors on
566	Prediction Skill as Estimated by Linear Inverse Modeling. <i>Journal of Climate</i> .
567	33(23), 10073–10095.
568	Roberts, C., Calvert, D., Dunstone, N., Hermanson, L., Palmer, M., & Smith, D.
569	(2016). On the drivers and predictability of seasonal-to-interannual variations
570	in regional sea level. Journal of Climate, 29(21), 7565–7585.
571	Rodgers, K., Lee, SS., Rosenbloom, N., Timmermann, A., Danabasoglu, G., Deser,
572	C others (2021). Ubiquity of human-induced changes in climate variabil-
573	ity. Earth System Dynamics, $12(4)$, $1393-1411$.
574	Rumelhart, D. E., Hinton, G. E., & Williams, R. J. (1986). Learning representations
575	by back-propagating errors. <i>Nature</i> , 323(6088), 533–536.
576	Sardeshmukh, P. D., & Sura, P. (2009). Reconciling non-gaussian climate statistics
577	with linear dynamics. Journal of Climate, 22(5), 1193–1207.
578	Schmid, P. J. (2010). Dynamic mode decomposition of numerical and experimental
579	data. Journal of Fluid Mechanics, 656, 5–28.
580	Shamekh, S., Lamb, K. D., Huang, Y., & Gentine, P. (2023). Implicit learning of
581	convective organization explains precipitation stochasticity. <i>Proceedings of the</i>
582	National Academy of Sciences, 120(20), e2216158120.
583	Shin, SI., & Newman, M. (2021). Seasonal predictability of global and north
584	american coastal sea surface temperature and height anomalies. <i>Geophysical</i>
585	Research Letters, $48(10)$, e2020GL091886.
586	Smith, R., Jones, P., Briegleb, B., Bryan, F., Danabasoglu, G., Dennis, J., others
587	(2010). The Parallel Ocean Program (POP) Reference Manual. LAUR-01853,
588	141, 1–140.
589	Stephenson, D., Hannachi, A., & O'Neill, A. (2004). On the existence of multiple cli-
590	mate regimes. Quarterly Journal of the Royal Meteorological Society, 130(597),
591	583-605.
592	Tziperman, E., Zanna, L., & Penland, C. (2008). Nonnormal thermohaline cir-
593	culation dynamics in a coupled ocean–atmosphere gcm. Journal of Physical
594	Oceanography, 38(3), 588-604.
595	Vimont, D. J., Alexander, M. A., & Newman, M. (2014). Optimal growth of Cen-
596	tral and East Pacific ENSO events. Geophysical Research Letters, 41(11),
597	4027–4034.
598	Vinogradova, N. T., Ponte, R. M., & Stammer, D. (2007). Relation between sea level
599	and bottom pressure and the vertical dependence of oceanic variability. Geo-

600	physical Research Letters, $34(3)$.
601	von Storch, H., Bürger, G., Schnur, R., & von Storch, JS. (1995). Principal Oscilla-
602	tion Patterns: A Review. Journal of Climate, 377–400.
603	Webb, D. J. (2021). On the low western Pacific sea levels observed prior to strong
604	East Pacific El Niños. Ocean Science, 17(6), 1585–1604.
605	Yeung, E., Kundu, S., & Hodas, N. (2019). Learning Deep Neural Network Rep-
606	resentations for Koopman Operators of Nonlinear Dynamical Systems. In 2019
607	American Control Conference (ACC) (pp. 4832–4839).
608	Zanna, L. (2012). Forecast Skill and Predictability of Observed Atlantic Sea Surface
609	Temperatures. Journal of Climate, $25(14)$, 5047–5056.

Supporting Information for "Learning Propagators for Sea Surface Height Forecasts Using Koopman Autoencoders"

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8. Table S3: Reconstruction MSE for different dimensionality reduction techniques on North Pacific daily SSH using different dimensionalities

Introduction

Here, we describe methodological training details and analysis metrics used in this study (Text S1, Text S2), provide supplementary figures describing the validity of the propagators (Figure S1 and S2), show the SSH variability due to different components to give context for the performance differences between regions (Figure S3), and provide tables to quantify the reconstruction and prediction performance of the different dimensionality reduction and propagation techniques (Tables S1, S2, and S3).

:

The encoder and decoder of our Convolutional Autoencoder and Koopman Autoencoder are composed of convolutional "blocks," where each block consists of a convolutional layer equipped with ReLU activations followed by another convolutional layer with ReLU activations (Fukushima, 1969, 1980). The convolutional layers use a 3-by-3 filter with a stride of 1 and employ zero-padding to preserve the shape of the input fields. In the encoder, convolutional blocks are succeeded by max-pooling operations using a 2-by-2 kernel, whereas in the decoder, convolutional blocks are preceded by bilinear upsampling using a 2-by-2 kernel. We use an architecture somewhat similar to Oommen, Shukla, Goswami, Dingreville, and Karniadakis (2022), where the number of filters per block is decreased closer to the bottleneck. In the encoder, the first two convolutional blocks contain convolutional layers with 64 channels, the third block contains layers of 32 channels, and the fourth contains layers of 16 channels. The last convolutional layer is fully connected to the latent space encoding. The decoder essentially has the reverse structure of the encoder: the encoding is fully connected to a convolutional block employing layers of 16 channels, followed by a block with layers of 32 channels, and then two blocks of 64 channels. Additionally, the decoder applies a 1-by-1 convolution to the outputs of the last convolutional block in order to return values in the range $(-\infty, \infty)$.

We optimize the parameters of the networks using the Adam optimizer (Kingma & Ba, 2014) with batches of 64 samples and a fixed learning rate of 10^{-4} . For the Koopman autoencoders, L_2 regularization is applied over all network weights to mitigate overfitting. For the daily-averaged data, an L_2 weight of 10^{-3} is applied, whereas for the monthly-

averaged data, a higher regularization rate of 10^{-2} was necessary to prevent overfitting. Networks are trained for 500 epochs, with an early stopping threshold of 50 epochs. Checkpoints for the network with the best overall validation loss were saved. Additionally, for the Koopman Autoencoder, we save the checkpoints with the best validation-set prediction MSE such that the learned propagator has decaying eigenvalues. This checkpoint with the best prediction loss is used.

The training capacity of both the Convolutional Autoencoder and Koopman Autoencoder was found to be sensitive to the network weight initializations: for certain initial weights, the network only converged to a constant function. Therefore, for the Convolutional Autoencoder, we initialize weights using Kaiming uniform random values (He et al., 2015), and reinitialize the weights with a different set of Kaiming uniform random values if the network does not converge to a lower loss than that of a constant function. For the Koopman Autoencoder, we leverage information gained about the loss landscape during the training process for the Convolutional Autoencoder. The Koopman Autoencoder's encoder and decoder weights are initialized from the weights of the Convolutional Autoencoder at the 10th epoch of training. This is based on the principle that lower-order features are learned first during training (Kalimeris et al., 2019; Refinetti et al., 2023): by beginning the training from the 10th epoch, the encoder and decoder contain enough complexity to converge to something more expressive than a constant function, but not so much complexity that the KAE overfits. Furthermore, the weights for the linear propagator L are initialized as a multiple of the identity matrix α **I**, where $\alpha \in (0, 1)$. Thus,

The data consists of 32,060 training samples for the daily data, and 21,014 samples for the monthly data (daily data is subsampled by a factor of 20 to reduce the computational cost). We use k = 20 recurrent passes for the prediction loss, and set the relative weights of the three different loss functions $\lambda_1 = \lambda_2 = \lambda_3 = 1$. The networks are trained in Pytorch using the distributed data parallel approach on two NVIDIA 32GB V100 GPUs (Paszke et al., 2019; Li et al., 2020).

Text S2. Metrics

Here we define metrics used for assessing reconstruction and prediction performance.

Let **X** be the tensor of target values for a specific geophysical field, and let $\hat{\mathbf{X}}$ be the predicted values. These tensors have entries $x_{i,j,n}$, and $\hat{x}_{i,j,n}$, where $i \in \{1, \ldots, M_x\}$ indexes the longitudes, $j \in \{1, \ldots, M_y\}$ indexes the latitudes, and $n \in \{1, \ldots, N\}$ indexes the samples.

We first define domain averaged metrics for a specific sample. Using a wildcard "*" to indicate dimensions of aggregation, the area-weighted Mean Squared Error (MSE) for a specific sample is given by

$$MSE_{(*,*,n)} = \frac{\sum_{i=1}^{M_x} \sum_{j=1}^{M_y} w_{i,j}^2 (x_{i,j,n} - \hat{x}_{i,j,n})^2}{\sum_{i=1}^{M_x} \sum_{j=1}^{M_y} w_{i,j}^2}$$
(1)

where $w_{i,j}$ gives the $(i,j)^{\text{th}}$ weight, which is proportional to grid-cell area on nondegenerate points and 0 on masked points. Similarly, the area-weighted pattern Correlation

Coefficient (CC) for a given sample is given by

$$CC_{(*,*,n)} = \frac{\sum_{i=1}^{M_x} \sum_{j=1}^{M_y} w_{i,j}^2 x_{i,j,n} \hat{x}_{i,j,n}}{\sqrt{\sum_{i=1}^{M_x} \sum_{j=1}^{M_y} (w_{i,j} x_{i,j,n})^2 \sum_{i=1}^{M_x} \sum_{j=1}^{M_y} (w_{i,j} \hat{x}_{i,j,n})^2}}$$
(2)

Global metrics over all gridpoints and samples can be obtained by averaging over all samples:

:

$$MSE = \frac{1}{N} \sum_{n=1}^{N} MSE_{(*,*,n)}$$
(3)

$$CC = \frac{1}{N} \sum_{n=1}^{N} CC_{(*,*,n)}$$
(4)

The area-weighted ℓ^2 -norms $\|\cdot\|_{2,w}$ given in Eqs. (3) and (4) use the globally-averaged area-weighted MSE in Eq. (3).

We can also consider the sample averaged MSE at each location, given by

$$MSE_{(i,j,*)} = \frac{1}{N} \sum_{n=1}^{N} (x_{i,j,n} - \hat{x}_{i,j,n})^2$$
(5)

It is often useful to assess the predictions of a model relative to another baseline. The skill score is an often used metric that assigns a value between 0 and 1 to assess the performance of the model relative to a baseline (Murphy, 1988). For a prediction model f and a baseline f_0 , we define the total skill score by

$$SS = 1 - \frac{MSE(f)}{MSE(f_0)}.$$
(6)

where MSE(f) gives the error given by the model f. This can be interpreted as the percentage of improvement in MSE gained by using model f instead of f_0 .

Likewise, the sample-averaged skill score for each location by

$$SS_{i,j} = 1 - \frac{MSE_{(i,j,*)}(f)}{MSE_{(i,j,*)}(f_0)}$$
(7)

where $MSE_{(i,j,*)}(f)$ is the sample-averaged MSE using prediction model f. Finally, domain-averaged skill is found by area-weighted averaging over all spatial indices (i, j):

$$\overline{SS} = \frac{\sum_{i=1}^{M_x} \sum_{j=1}^{M_y} w_{i,j} SS_{i,j}}{\sum_{i=1}^{M_x} \sum_{j=1}^{M_y} w_{i,j}}$$
(8)

References

- Fukushima, K. (1969). Visual Feature Extraction by a Multilayered Network of Analog Threshold Elements. *IEEE Transactions on Systems Science and Cybernetics*, 5(4), 322–333.
- Fukushima, K. (1980). Neocognitron: A Self-organizing Neural Network Model for a Mechanism of Pattern Recognition Unaffected by Shift in Position. *Biological Cybernetics*, 36(4), 193–202.
- He, K., Zhang, X., Ren, S., & Sun, J. (2015, December). Delving Deep into Rectifiers: Surpassing Human-Level Performance on ImageNet Classification. In Proceedings of the IEEE International Conference on Computer Vision (ICCV).
- Kalimeris, D., Kaplun, G., Nakkiran, P., Edelman, B., Yang, T., Barak, B., & Zhang,
 H. (2019). SGD on Neural Networks Learns Functions of Increasing Complexity. In Advances in Neural Information Processing Systems (Vol. 32). NeurIPS.
- Kingma, D. P., & Ba, J. (2014). Adam: A method for stochastic optimization. arXiv preprint arXiv:1412.6980. doi: 10.48550/arXiv.1412.6980
- Li, S., Zhao, Y., Varma, R., Salpekar, O., Noordhuis, P., Li, T., ... Chintala, S. (2020, aug). PyTorch distributed: experiences on accelerating data parallel training. *Proceedings of the VLDB Endowment*, 13(12), 3005–3018. Retrieved from https://doi.org/10.14778/3415478.3415530 doi: 10.14778/3415478.3415530

Murphy, A. H. (1988). Skill Scores Based on the Mean Square Error and their Relationships to the Correlation Coefficient. *Monthly Weather Review*, 116(12), 2417–2424.

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- Oommen, V., Shukla, K., Goswami, S., Dingreville, R., & Karniadakis, G. E. (2022). Learning two-phase microstructure evolution using neural operators and autoencoder architectures. npj Computational Materials, 8(1), 190.
- Paszke, A., Gross, S., Massa, F., Lerer, A., Bradbury, J., Chanan, G., ... others (2019). Pytorch: An Imperative Style, High-Performance Deep Learning Library. Advances in Neural Information Processing Systems, 32.
- Penland, C., & Sardeshmukh, P. D. (1995). The optimal growth of tropical sea surface temperature anomalies. *Journal of Climate*, 8(8), 1999–2024.
- Refinetti, M., Ingrosso, A., & Goldt, S. (2023). Neural networks trained with SGD learn distributions of increasing complexity. In *International Conference on Machine Learning* (pp. 28843–28863).



Figure S1. Eigenvalues of discrete propagators of LIM, $\mathbf{B}(1)$, for both PCA and CAE latent modes, as well as the eigenvalues of the Koopman Autoencoder propagator L. The unit circle demarcates the region in which the eigenvalues must lie for the propagator to be stable.

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Figure S2. PCA+LIM evolution matrix norms by fitted propagator lead time. The blue line shows the matrix norm itself, with a star indicating the model with the lowest average prediction MSE over timesteps 1-k on the validation dataset. The red line shows the norm of an average propagated latent space vector σ , as in Penland and Sardeshmukh (1995) Fig. 12.



Figure S3. Explained variance of daily SSH variability by component in the North Pacific (a, b, c) and North Atlantic (d, e, f). Panels (a) and (d) show the proportion of SSH variability due to dynamic sea level, while panels (b) and (e) show the proportion due to the inverse barometer component. Because the random variates ζ and η_{ib} are not completely decorrelated, the explained variance by the two terms do not exactly sum to 1. Therefore, the closure term due to covariance $2\text{Cov}(\zeta, \eta_{ib})/\text{Var}(\eta)$, which is negligible at most locations, is included in panels (c) and (f).

Table S1.	Reconstruction	MSE for	different	dimension	ality	reduction	techniques	with a	d = 2	20
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	Pacific Daily	Atlantic Daily	Pacific Monthly	Atlantic Monthly
model				
PCA	0.191	0.065	0.167	0.082
CAE	0.185 (-3.61%)	0.063 (-2.00%)	0.161 (-3.83%)	0.086 (+5.26%)
KAE	0.198 (+3.27%)	0.078 (+20.69%)	0.231 (+38.39%)	0.135 (+64.10%)

latent dimensions. Parentheses show the percent difference in MSE from PCA.

 Table S2.
 Total skill score (expressed as a percentage) of different prediction methods relative

to PCA-DP, averaged over forecast leads up to 120 days for daily data and 36 months for monthly

data.

	Pacific Daily	Atlantic Daily	Pacific Monthly	Atlantic Monthly
Prediction method				
CAE-DP	-15.4%	-0.4%	-1.0%	-0.9%
PCA-LIM	9.4%	0.6%	13.9%	3.3%
CAE-LIM	9.6%	0.2%	12.8%	3.0%
KAE	13.9%	1.7%	16.0%	2.0%

 Table S3.
 Reconstruction MSE for different dimensionality reduction techniques in the North

Pacific on daily timescales for different numbers of latent dimensions. Lighter shading represents

lower MSE. Parentheses show the	percent difference in MSE from PCA.
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	D=10	D=20	D=30	D=40
Technique				
PCA	0.308	0.191	0.137	0.106
CAE	0.301 (-2.35%)	0.185 (-3.61%)	0.131 (-4.01%)	0.103 (-2.11%)
KAE	0.311 (+0.92%)	0.198 (+3.27%)	0.149 (+8.80%)	0.119 (+12.72%)